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# Groundwater Resource in the Crystalline Rocks of the Johannesburg Area, South Africa

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## Abstract

Understanding the groundwater dynamics is very important for strategic management of the water resource as urbanization and population growth put South Africa's current water supply under tremendous stress. The groundwater resources potential and quality characteristics of the greater Johannesburg area, was thoroughly assessed using hydrogeochemical and environmental isotope techniques. The results of the geochemical and stable isotope investigations help understand the groundwater setting. Mixing process of fresh and polluted water is taking place at shallow zones within the weathered crystalline rocks and dissolution cavities in dolomites, where the rocks are characterized by the lithological and structural complexities. Furthermore, it was confirmed that the solute composition of the water results from two basic contributions: the reaction of meteoric water with weathered zones in the near surface area and direct infiltration through dolomite dissolution cavities. The results further indicate that the dominant movement mechanism of ions is through a diffusive process, which could be due to the limited availability of fractures with depth and its possible sealing by calcite precipitation. Consequently, the provenance of groundwater circulation is limited to shallow depths, however, dissolution cavities, fractured and sheared zones allow deeper circulation of groundwater. Moreover, the natural water quality is found to be highly impacted by acid mine drainage which is derived from the gold mines in the Witwatersrand basin as it has been observed in previous studies.

**Keywords:** Crystalline Aquifer, Diffusive Circulation, Dissolution Cavities, Johannesburg, Water Supply

## 1. Introduction

Most countries, located within the arid and semi arid regions like South Africa, are categorized as water stressed [1,2] and depend primarily on groundwater as water supply for various developmental activities. Due to large water usage, the main water supply for big cities like Johannesburg and Pretoria is derived from surface water sources. However, many rural communities rely on groundwater for domestic supply, which is obtained from shallow aquifers composed of alluvials and weathered crystalline rocks. Groundwater also plays strategic role in supplying water for large irrigation fields in the area which is extracted from deep wells located within the dolomitic aquifer in the Transvaal Supergroup [3-5]. Therefore, understanding the nature of groundwater circulation and hydrogeochemical characteristics of the aquifers is crucial for sustainable development of the coun-

try's water resources. Previous knowledge on the groundwater circulation media, natural quality of water could give insight on the state of pollution, which are generated from urban, irrigation, mining and industrial activities.

**Figure 1** illustrates location of the study area which is characterized by semi arid climate with the mean annual rainfall ranging from 600 to 700 mm/year [6], showing high variability with the majority of rain falling during summer months (December through February). The summer rainfall in the area is extremely important in its contribution for groundwater recharge.

Due to the presence of complex and heterogeneous crystalline aquifers in the Johannesburg area, the identification of productive aquifers has become problematic and recharge is presumed to be small due to the crystalline nature of the rocks. Past studies clearly show that major ions and environmental isotopes are proven to be

important in determining the characteristics of groundwater in complex aquifer systems [7]. The integrated application of the methods has also been successfully used in the characterization of the geochemical evolution in granitic batholiths in Canada and other parts of the world [7-9]. Because of the spatial variability in the degree of fracturing and weathering in the crystalline rocks, the development of water bearing horizons is limited to shallow weathering zones and fracture media. On the other hand, the dolomitic aquifers are identified as highly productive aquifers in the region [5,10].

The study area and the adjoining Witwatersrand basin are well known for its large-scale urbanization and industrial activities with substantial water demand. The region hosts one of the world's largest gold and uranium mining basins having more than 120 mines. In the mining areas, deeper groundwater was pumped out at large scale with a rate of 110 Mega liters per day (ML/day) until 2008 to make underground mining possible [11]. However, evidences confirm that the deep underground mine pumping does not tremendously affect the shallow groundwater. Surface discharge of largely polluted pumped water to dams, reservoirs and local streams considerably increase recharge of the shallow aquifer [12].

According to [13] the Cradle of Human Kind World Heritage site which is located in the western part of Johannesburg is seriously threatened by the acid mine drainage. Ground and airborne geophysical surveys have also identified sub surface acid mine drainage pathways in Krugersdorp Game Reserve area [14]. Recent estimates of the rate of decant ranges between 18 ML/d and 36 ML/day with a pH  $\approx$  3 and electrical conductivity of  $\approx$  550 mS/m [15]. In the area, groundwater is highly used for domestic and industrial activities. With growing fresh water scarcity; groundwater resource assumes increasing importance and needs protection according to best management practice. Hence, it is important to understand the groundwater recharge and circulation mechanisms and quality in these crystalline rocks. Proper understanding of the geochemical process will further help in managing the groundwater quality in the region.

The elevation in the area ranges from 1 365 meters above sea level (m a.s.l.) at Pretoria to 1 753 m a.s.l. at Johannesburg (with an elevation difference of 388m). The central section of the city of Johannesburg is situated on the major water divide between the northern Limpopo River basin with numerous networks of tributaries and the southern Orange River basin (see **Figure 1**).

For this work, hydrogeochemical and isotope variables were considered as they give important indications regarding the geoenvironmental history of the enclosing rocks, the velocity and direction of movement and the mechanism of flow in the aquifers. The spatial variation

in the concentration of dissolved substances determines the chemical properties and evolution of water along a flow path. In the study area, which is covered with large variety of basement lithologies, the large variation in water chemistry reflects the degree of fracturing and weathering of the host rocks.

## 2. Objectives and Methods

The main purpose of this work is to assess the quality variation, aquifer characteristics and hydrogeochemical nature of groundwater within the basement crystalline rocks of Johannesburg area. In order to understand the groundwater dynamics, groundwater database from the Department of Water and Environmental Affairs (DWEA), Pretoria, has been utilized with field verification on accessible water points for physical parameters such as pH, Electrical Conductivity (EC) and Total Dissolved Solids (TDS). Additional samples were collected in 2009 and 2010 from western part of Johannesburg for environmental isotope analysis. Field visits were conducted in order to understand the lithostratigraphic relationship which helped for the preparation of the simplified geological map. There is an ongoing surface and ground water interaction project in the area and this paper reports the first results.

## 3. Geological Framework

The Johannesburg area is composed of crystalline rocks of Archean age. They are broadly classified as granitic rocks, meta-sedimentary and meta-volcanics (**Figure 2**). The Johannesburg dome which is located between Johannesburg and Pretoria which is circular in outline consists of Archean greenstones remnants (c. 3.34 Ga, [16] and intruded by tonalite, granodiorite, granite and migmatites, unconformably overlain by metasedimentary rocks of < 3 000 Ma [17-19]. The tonalites are the oldest rock types, migmatites are intermediate in age and granodiorite and granite are the youngest [19]. The oldest granitic rocks comprise a suite of tonalitic and trondjemitic gneisses and migmatites that occupy most of the northern half of the dome. Exposure of similar rocks also occurs on the southern edge of the dome and unconformably underlies the Witwatersrand Supergroup [19], [20].

The Witwatersrand basin is an arcuate structural basin lying within the Kaapvaal Craton. The Witwatersrand Supergroup is divided into the Lower West Rand Group and the upper Central Rand Group. The former conformably overlies volcanics of the Dominion Group and non-conformably overlaps Archean basement rocks of the Kaapvaal Craton. The south-central portion consists

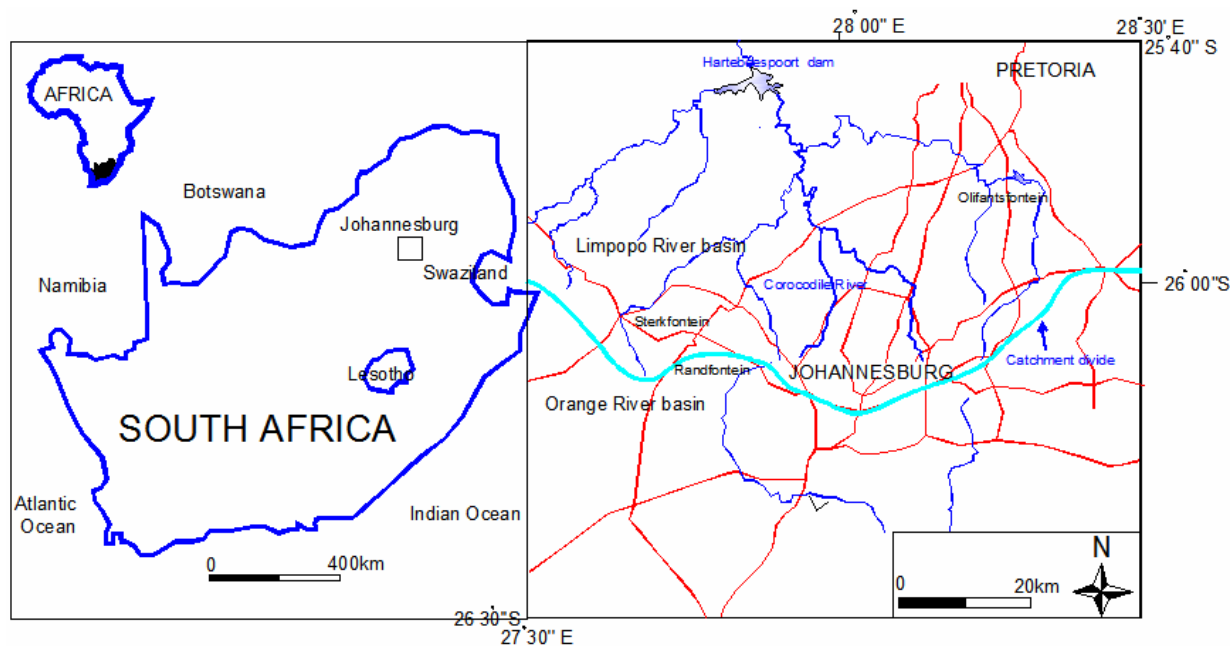


Figure 1. Location map of the study area, Johannesburg at the north central part of the country (east-west solid line is water divide between Limpopo and Orange Rivers, blue lines represent streams and red lines represent roads).

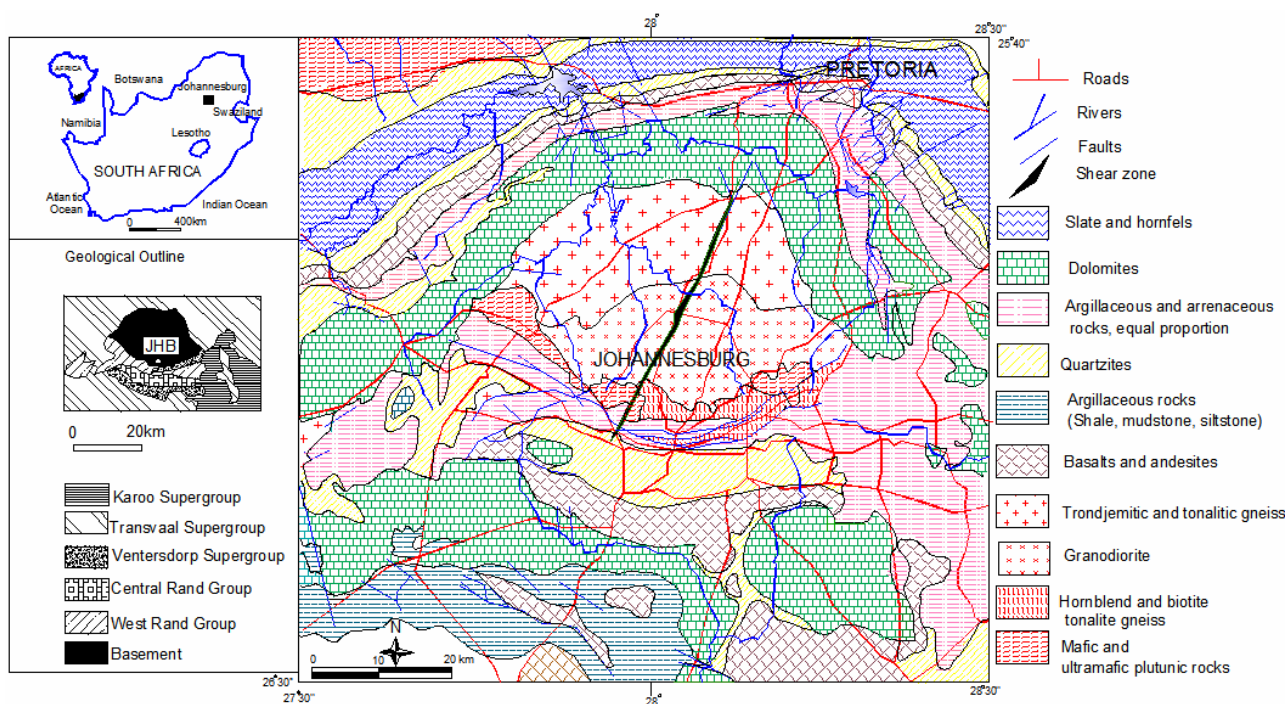


Figure 2. Simplified geological map of the Johannesburg area (Compiled from [17,25,27]).

mainly of a variety of homogeneous, medium grained granodioritic rocks. Locally, all rocks contain a gneissic fabric and are less developed in the granodiorite and the tonalities. The Witwatersrand Supergroup unconformably overlies basement granitoids and greenstones, as well as sedimentary and volcanic rocks of the Dominion

Group. It is made of thick terrigenous sequence comprising arenaceous and argillaceous sedimentary rocks. The deposition took place between 3074 and 2 714 Ma [21]. The part of Witwatersrand basin that lies close to Johannesburg dome is grouped under West Rand, Central Rand and East Rand Groups. The west rand group con-

sists primarily of quartzites and shales. The Central Rand Group consists of different proportion of quartzites and shales where the sequence consists mainly of quartzites and conglomerates [22].

The Early Proterozoic Transvaal Supergroup comprises of relatively undeformed, unmetamorphosed volcanic rocks, quartzites, shales, dolomites iron formations, conglomerates and diamictites [23]. The Transvaal Supergroup is made of relatively undeformed, unmetamorphosed volcanic rocks, quartzites, shales, dolomites, iron formations, conglomerates and diamictites [23]. The Supergroup includes dolomites and iron formations of Chuniespoort Group [23] which is overlain by Pretoria Group that comprises of predominant mud rocks alternating with quartzitic sandstones, significant interbedded basaltic-andesitic lavas and subordinate conglomerates, diamictites and carbonate rocks all of which have been subjected to low grade metamorphism [24].

#### 4. Hydrogeological Framework

The present landscape of the study area, which is marked by elongated ridges, rolling topography and wide plain areas could be a clear reflection of the old geological activity (intrusion, sedimentation, metamorphism, ductile and brittle tectonics) and subsequent erosion processes. These activities have tremendous impact on hydrogeological characteristics of the rocks primarily by regulating recharge into the groundwater. The Johannesburg dome, having a circular outline, form slopping relief terrain whereas the enclosing younger rocks (Transvaal Sequence) form relatively high ground. Fractures of all orientation and weathering zones exist in all rocks. From hydrogeological point of view, the rocks that outcrop in the study area fall under hardrock category with low groundwater productivity except dolomites that contain dissolution cavities and consequently host huge quantity of groundwater.

According to the Johannesburg hydrogeological map of 1: 500 000 scale [25], and explanatory note [10]; four aquifer types have been identified in the area. Namely:

- An intergranular aquifer in the alluvial covered zones
- Fractured aquifer in the Witwatersrand Supergroup associated with fractures, fissures and joints
- The karstic aquifer in the dolomites
- The intergranular and fractured aquifer in the crystalline rocks

However, from detailed field surveys, it was possible to conceptualize that groundwater occurrence in the area can be categorized into three broad groups: near surface occurrence within the weathered profile; occurrence within fractures, dykes and shear zones; and occurrence

within dissolution cavities in the dolomites. The main characteristic features of groundwater in the highly productive dolomitic aquifer [3-5] are pockets of conduits which are compartmentalized primarily by structural discontinuities such as left lateral strike slip faults. The dolomitic formations generate little surface run-off, suggesting relatively high recharge and predominance of underground water flow, which eventually drains to springs typically associated with dykes, faults or formation contacts [13]. However, it has been observed that dikes, sills and quartz veins generate local perched system which frequently generates springs in the area. The presence of recharging meteoric water through dykes and faults and the influx of thermal water along fractures within the deep gold mines in the Witwatersrand basin have been documented [26]. Therefore, occasionally dykes and faults play important role in regulating recharge and, channeled and deep groundwater circulation zones in the area.

The semicircular deformation of the Witwatersrand and Transvaal Supergroup rocks around Archean granitic pluton (see the central part in **Figure 2**) could play substantial role in regulating groundwater circulation in the area. In the same area, the randomly oriented left lateral strike slip faults with minor occurrence of right lateral strike slip faults with the associated weathered and sheared zones can potentially act as a conduit for accelerated groundwater flow. Since dikes constitute an important structural set up of the crystalline rocks in the area, they act as conduits for groundwater circulation besides compartmentalizing role in some instances. Owing to the fact that the area is made of deformed crystalline rocks, the main hydrostructures are represented mainly by fractures and weathered zones. The fractured Black Reef quartzites of the Witwatersrand Group that form the water divide between the northern Limpopo River basin and the southern Orange River basin are intensively fractured, which provide suitable media for vertical recharge. Granitic gneisses and quartzites, instead, are characterized by massive structures with wide fracture network while shales contain very tight fractures along inclined beds to the south. The conceptual model demonstrates that the city of Johannesburg is located on the water divide where groundwater flows away towards north and South (**Figure 3**). However, the hydraulic connection exists between different lithological units. The Transvaal dolomites in the western and eastern part of Johannesburg area, groundwater moves across the water divide through dissolution cavities and tectonic lineaments (**Figure 2**). Alluvial deposits found along several stream valleys in the area are highly productive, such as aquifers in the the lower Crocodile River valley (down-stream of Hartbeespoort dam), with

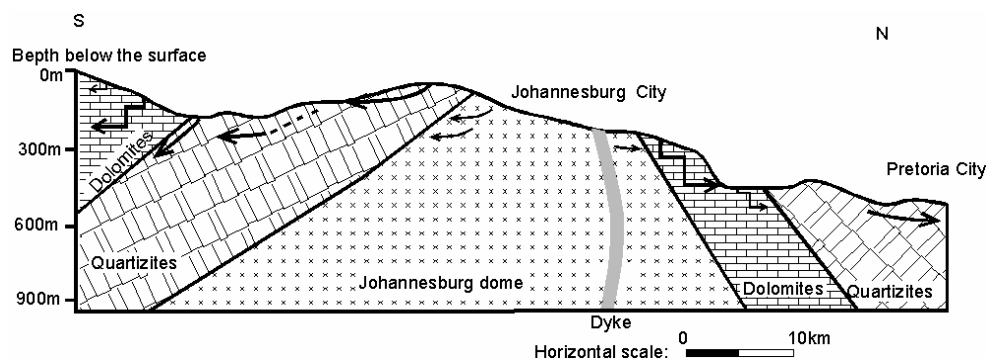


Figure 3. Conceptual hydrogeological model along south-north section in Johannesburg.

Table 1. Chemical composition of some water supply wells in the Johannesburg area.

Code	Latitude	Longitude	Elev. m.a.s.l	SWL (m)	Well Depth (m)	pH	EC mS/cm	K mg/l	Na mg/l	Ca mg/l	Mg mg/l	Cl mg/l	TAL mg/l	SO <sub>4</sub> mg/l
Well_1	-26.036 12	27.7	1520	7.92	9.14	7.7	112.5	1.36	62.5	93.3	53.3	94.3	97.6	227.1
Well_2	-25.856 39	28.166 95	1430	15	31	7.6	66	1.63	6.8	69.6	38.9	15.4	263.3	20.4
Well_3	-25.896 39	28.312 5	1487.83	22	38	7.7	53.9	0.54	16.1	45.4	31.1	20	172.4	32.1
Well_4	-25.927 5	28.219 45	1480	30	45	7.28	87.6	1.76	61.3	63.6	30.1	78.2	209.7	59.1
Well_5	-26.017 4	27.711 36	1463.52	30	56	7.2	98	1.51	46.1	87.8	49.1	70.2	119.8	229.7
Well_6	-26.080 56	27.576 39	1545.22	53	60	6.8	83	0.7	1	7.1	4	1.5	42.4	9.9
Well_7	-25.898 62	28.302 78	1519.46	67	69	7.4	37.1	1.38	5.8	35.2	20.6	6.6	121.3	21.6
Well_8	-25.969 45	28.280 56	1586.23	76	83	7.75	40.1	0.43	4.9	39.4	23.9	6.4	168.8	11.4
Well_9	-25.916 89	28.338 56	1500	10	87	7.85	24.6	0.15	1	23.8	12.8	1.5	118.7	5.5
Well_10	-25.889 16	28.208 88	1460	40	89	7.4	56	0.92	9.3	46	37.7	18.1	224.2	7
Well_11	-25.892 22	28.321 95	1524.21	63	92	7.58	255	0.62	5.6	24.2	15	4.6	92.3	8.1
Well_12	-26.068 25	27.647 36	1573.01	92	98	7.64	373	0.54	14.2	30.5	19.8	26	115.2	15.8
Well_13	-26.050 8	27.672 48	1546.77	74	108.5	7.66	34.4	1.6	7.6	35.1	21.2	39	98.5	47.3
Well_13	-25.882 22	28.256 66	1460	107	114	7.2	44.8	0.65	11	41.4	25.9	11	192.5	12.6
Well_14	-26.016 67	28.285 55	1570	116	117	7.47	50.2	1.28	12.7	39.8	31.4	28.5	158.6	33
Well_15	-25.908 61	28.316 95	1501.21	134	135	7.8	64	0.86	51	39.9	24.1	41.3	153.2	76.1
Well_16	-26.046 25	27.678 72	1543.61	132	135	7.82	31.6	0.4	3.2	30.8	19.3	10.9	125.8	8.3
Well_17	-25.85694	28.233 33	1490	148	158	7.65	60.4	2.73	6.4	62.3	35	3	265.9	15.6
Well_18	-25.949 61	28.341 94	1580	19	164	7.48	20.4	0.38	1	21.3	11	1.5	106	6
Well_19	-25.95	28.268 06	1559.77	169	176	7.7	36	0.79	5.4	36.1	21.6	4.2	157.3	9.8

blow yield as much as 16 l/s [10]. The dolomite areas are traversed by impervious and semi-pervious syenite and diabase dykes, which divide these areas into separate groundwater compartments [14]. In the dolomitic aquifer, the occurrence of cavities varies from place to place which is based on the intercalations found in dolomites.

Such variability depends on different generation of dolomite (dolomite containing chert layers-**Eccles Formation** and those without chert-**Frisco Formation**), the availability of acidic infiltrating water, the amount of groundwater moving through the system and the degree

of saturation with respect to calcite and dolomite. In large part of the area, dolomites have notorious reputation of developing sinkholes due to the formation of dissolution cavities [14]. The presence of acid mine drainage which is generated from the slime dams and abandoned shafts from gold mines have a tremendous potential to enhance dissolution of dolomites in the area. The occurrence of a series of dissolution cavities along tectonic lines is a characteristic feature in the dolomites. Appearance and disappearance of springs is also a common feature in the area, which is controlled by availability

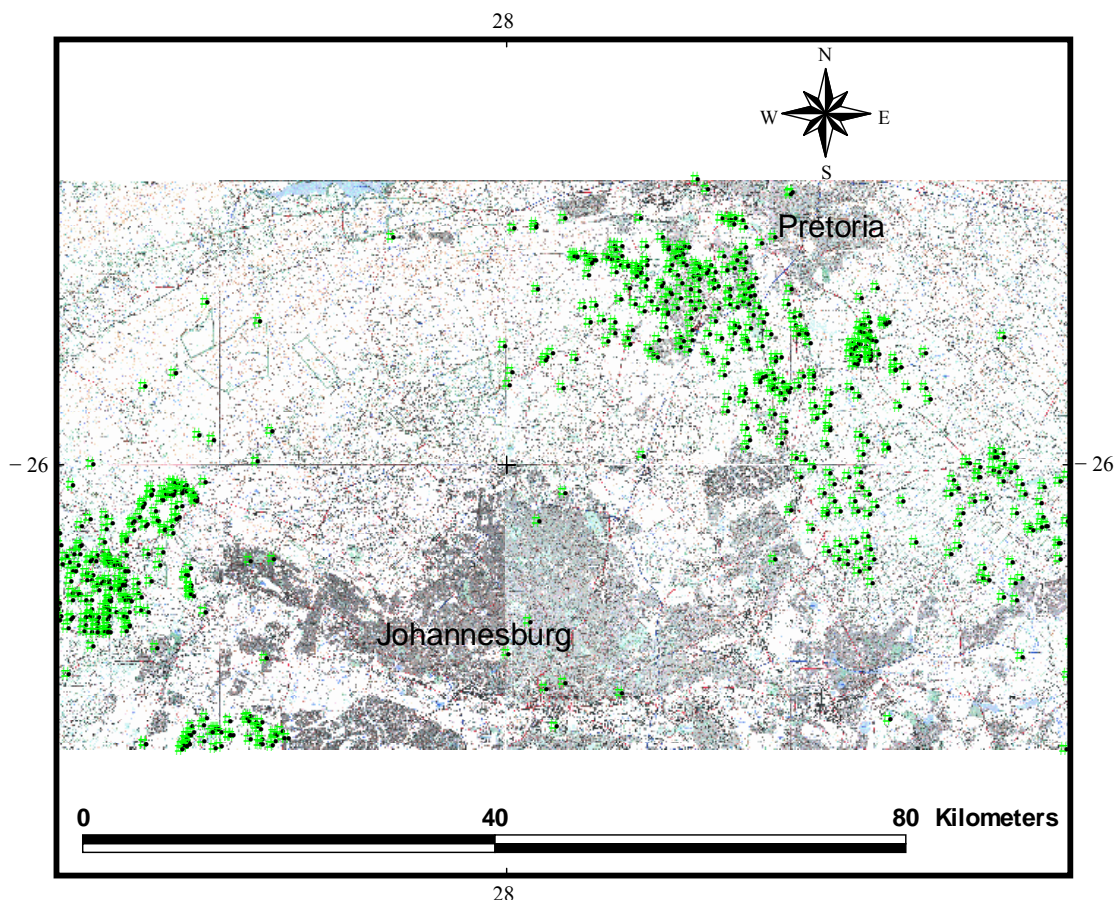


Figure 4. Location map for some water supply wells around Johannesburg.

of karst structures.

From 722 borehole data obtained from DWEA, Pretoria, it was possible to observe variation in borehole yield. The data have been interpreted in such a way that poorly drained crystalline aquifers produce from 0.01 L/s to 0.98 L/s, while weathered and fractured aquifers produce between 1 L/s and 14.6 L/s. High discharge wells are tapping groundwater from dolomites with the rate that varies between 15 L/s and 124 L/s (DWEA database).

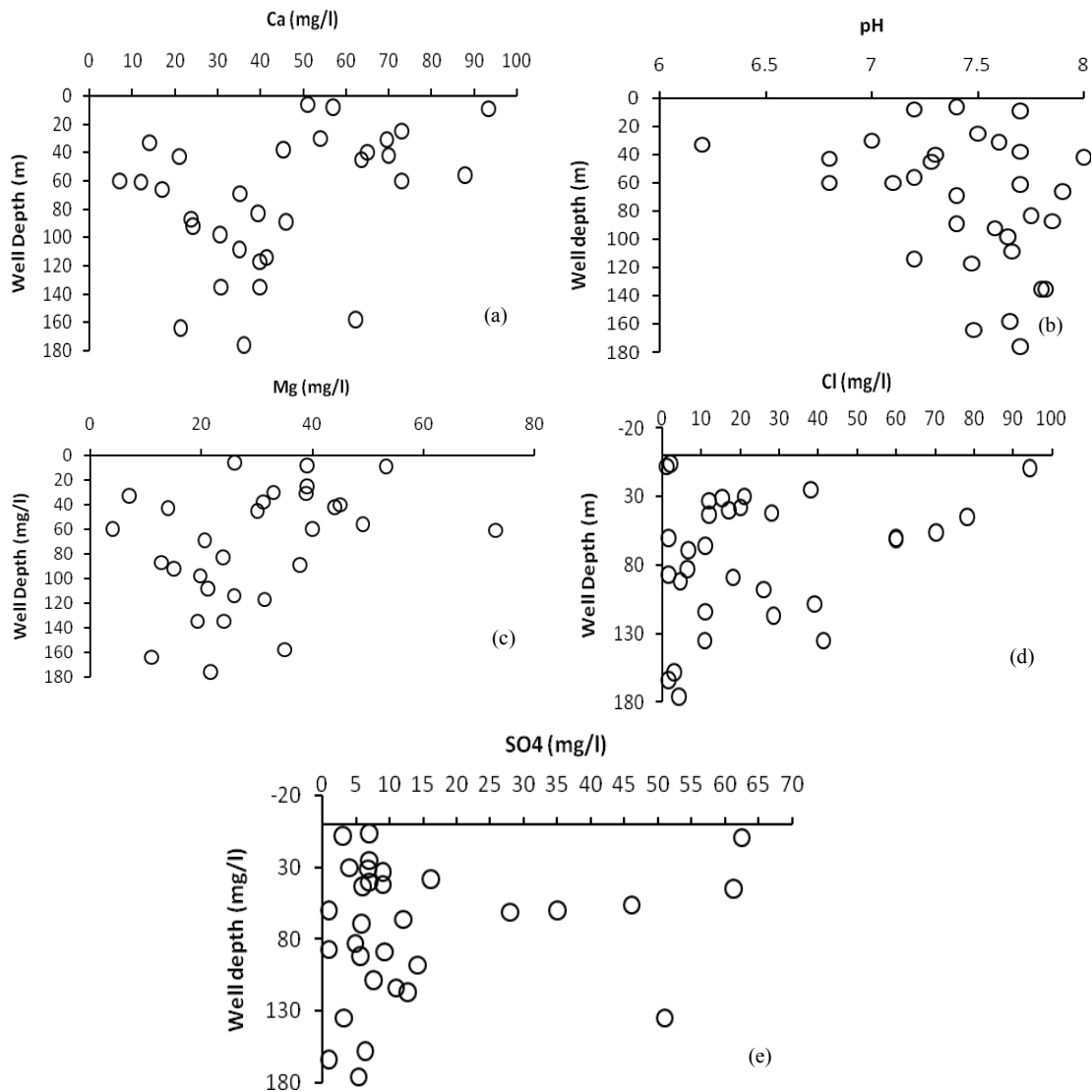
## 5. Results and Discussion

### 5.1. Physico-Chemical and Analytical Data

To assess the hydrogeochemical variation in the study area, different combinations of parameters have been plotted and some selected data have been presented in **Tables 1** and **2**. The distribution of water points (sampling points) is presented in **Figure 4**. The plots depict one prominent cluster which represent dominant aquifer in the area and dispersed points with high salinity that is related to pollution from gold mining areas. In different aquifers, groundwater contains high salinity due to leaching from

weathering profile and pollution derived from gold mines. For the selected boreholes, the distribution of ions with depth has been presented in **Figure 5** and the plots portray a relatively decreasing pattern with depth, which, instead, suggests diffusive circulation process due to calcite precipitation within fractures (**Figure 5(a)**). It is supported by the increasing pH (**Figure 5(b)**) due to carbonate saturation.

The increase in pH with depth (**Figure 5(b)**) enhances loss of calcium through precipitation process [8] and possible cation exchange with sodium. The lower pH values at shallow circulation zones could be facilitated by the presence of pollution from mines. In low or no flow condition, dissolved salts move mainly by diffusive rather than advective processes. The diffusive circulation could be attributed to the fracture sealing by precipitated calcite that reduces the available pore space at depth. At shallow depth the presence of relatively less alkaline water could be due to the influence of atmospheric and biogenic carbon dioxide through open fractures and dissolution cavities besides the pollution derived from acid mine decant. The decrease in the concentration of magnesium with depth (**Figure 5(c)**) reflects less solu-



**Figure 5. Distribution pattern of chemical constituent with the total well depth**

ble nature of dolomite with chert intercalation that results diffusive movement of ions in the dolomitic aquifer. High concentration of chloride and sulphate are evident, up to the depth of 60 m, likely due to the impact of pollution from mines (**Figures 5(d)** and **5(e)**). Chloride could be added to the aquifer from soluble salts from the weathered profile, rock water interaction (including reaction with matrix minerals, fracture-filling minerals) and the presence of soluble salts.

To evaluate the control pattern of various ions on EC, which has intimate relation with salinity, 2D variation diagrams have been prepared (**Figures 6(a)** to **(d)**). The plots show the presence of one dominant cluster (left lower corner) that has low salinity water and the other with highly conductive polluted water at the top right part of the graphs which are related to water from gold

mining areas. On same diagrams one can observe the fast enrichment of ions which is related to rock dominance phase with an additional source from multilayer aquifer system or mineral leaching. An interesting observation is evident regarding the fast increase of silica which is believed to be related to silicate dissolution from chert-rich dolomites and granites in more acidic and oxic conditions, which could be assisted by widespread Acid Mine Drainage (AMD) process in the area. The pattern on **Figure 6**, in general, indicates similar source through weathering of different rock types and mixing of different water types is also a possible controlling process for linear behavior of different water types. The distribution of points on the variation diagrams indicates the importance of all ions at low EC condition (lower left part). The points that lie in right top



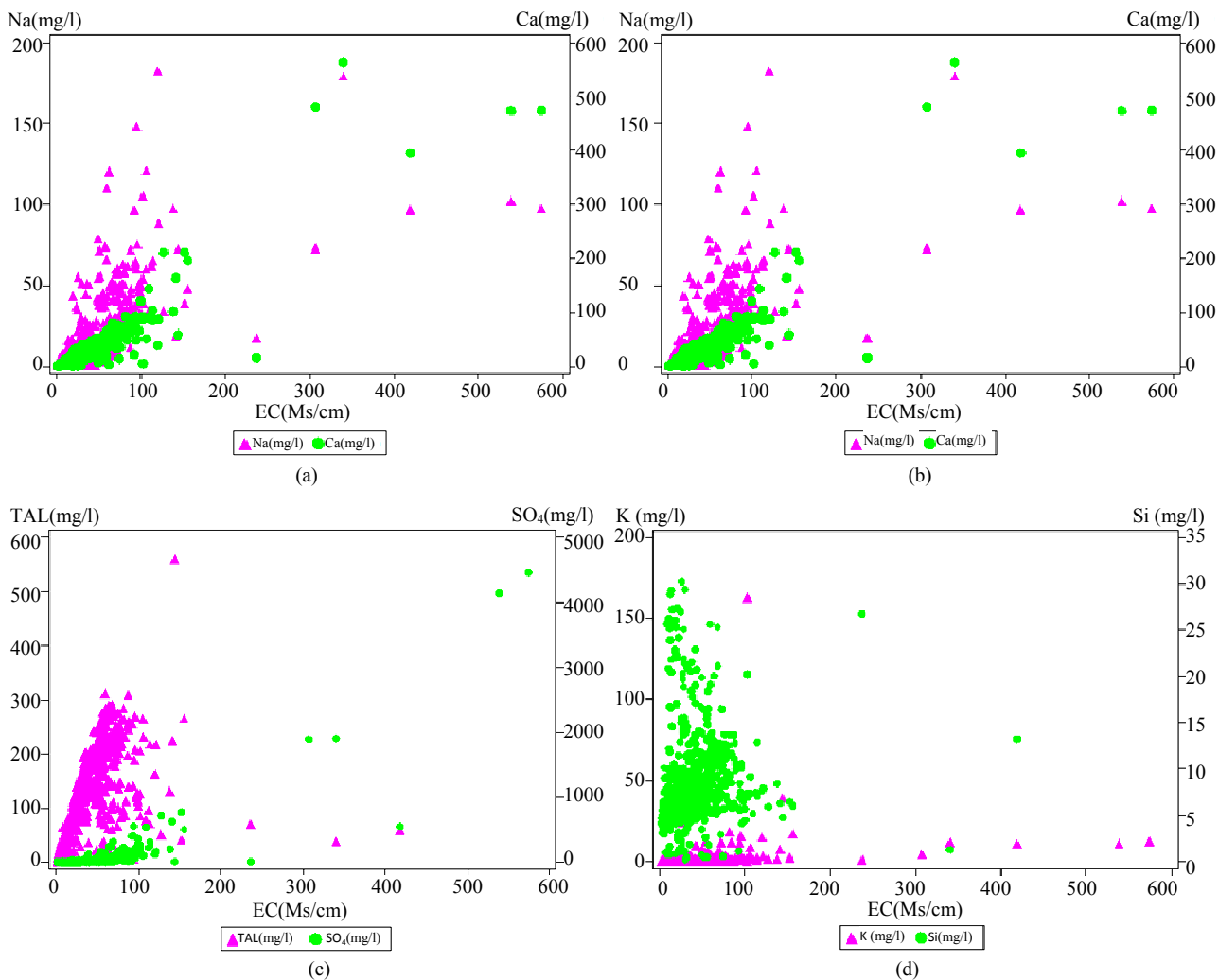


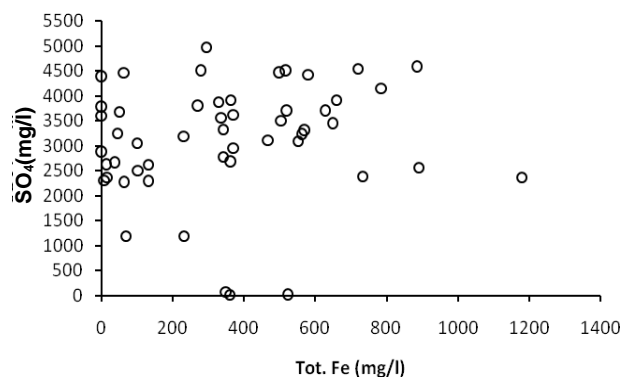
Figure 6. Control pattern of different ions over Electrical conductivity.



Figure 7. Acid mine decant pictured from one of the streams in the west rand gold mine areas

**Table 2. Concentration of total iron and sulphate in the groundwater which is affected by acid mine drainage.**

Sample code	Latitude	Longitude	SO <sub>4</sub> mg/l	Fe (tol) mg/l
W2	-26.1081	27.722 78	3238.15	563.01
W3	-26.1081	27.722 78	1194.37	231.77
W4	-26.1081	27.722 78	1194.37	69.375
W5	-26.1081	27.722 78	2620.67	132.78
W6	-26.1081	27.722 78	2692.11	361.8
W9	-26.1081	27.722 78	2372.03	1179
W10	-26.1081	27.722 78	2560.66	890.4
W14	-26.1081	27.722 78	2629.64	14.313
W15	-26.1081	27.722 78	2505.04	102.1
W16	-26.1081	27.722 78	2294.46	132
W18	-26.1136	27.722 78	4470.65	498.08
W19	-26.1136	27.722 78	4514.29	516.49
W21	-26.1136	27.722 78	4423.07	579.24
W22	-26.1136	27.722 78	2390.39	732.46
W23	-26.1136	27.722 78	4592.93	884.71
W24	-26.1136	27.722 78	4546.96	720
W26	-26.1136	27.722 78	4512.53	279.28
W27	-26.1136	27.722 78	3502.78	503
W28	-26.1136	27.722 78	3457.99	649
W29	-26.1136	27.722 78	3621.41	369.66
W30	-26.1136	27.722 78	3810.06	270.04
W32	-26.1136	27.722 78	3677.8	51.45
W33	-26.1136	27.722 78	3567.2	335.65
W34	-26.1136	27.722 78	3318.49	568.7
W35	-26.1136	27.722 78	4144.95	783.7
W36	-26.1136	27.722 78	3914.46	659.1
W37	-26.1136	27.722 78	3710.56	627.7
W40	-26.1133	27.722 78	3709.5	519.32
W41	-26.1133	27.722 78	2953.56	369.81
W43	-26.1133	27.722 78	3917.12	362.65
W44	-26.1133	27.722 78	3878.4	329.25
W45	-26.1133	27.722 78	3324.1	342.5
W46	-26.1133	27.722 78	3116.56	467.1
W47	-26.1133	27.722 78	4974.11	294.95
W48	-26.1133	27.722 78	2778.31	341.8
W51	-26.1133	27.722 78	3186.21	230.86
W52	-26.1133	27.722 78	3087.93	551.9
W53	-26.1133	27.722 78	2961.63	492.7
W55	-26.1153	27.725	4176.01	83.69
W58	-26.1153	27.725	2644.78	986.8

**Figure 8. The distribution of Total Iron and SO<sub>4</sub> within the boreholes in Johannesburg gold mine area.**

part of **Figure 6** are related to AMD impact in the shallow boreholes.

Total Alkalinity (TAL) increases rapidly (**Figure 6(c)**) at an initial stage indicating the presence of multiple sources (carbonate weathering and CO<sub>2</sub> reaction in the soil). The influence of atmospheric and soil CO<sub>2</sub> probably accounts for increase in TAL values primarily in shallow aquifer. Due to calcite precipitation with depth, TAL is expected to decrease further below 200m depth. Similar enrichment patterns of ions indicate mixing of different water and can be considered as an important process to generate these ions.

The impact of Acid Mine Drainage (AMD) is the most dominant environmental problem in the area due to its capacity to dissolve carbonates and direct outflow from mine shafts into the streams as pink coloured flow (**Figure 7**) with high concentration of total iron and sulphate (**Table 2** and **Figure 8**). The data presented in **Table 1** confirm that the groundwater system is influenced by AMD which facilitates mineral dissolution process dominantly in the shallow aquifer system.

## 5.2. Principal Component Analysis

Principal Component Analysis (PCA) a statistical tool used to discriminate common categories out of a data cluster to draw sensible and logical conclusion. In previous investigations, PCA was commonly applied for discriminating different water groups in various parts of the world such as in [26,28-30], etc. Therefore, in the current work PCA is used to identify chemistry of dominant groundwater types thereby identifying common geochemical categories and grouping of water quality types besides the interrelations to the aquifer can be done. For the PCA, seven dominant ions/parameters were considered; namely Ca, Mg, Na, K, Cl, SO<sub>4</sub>, and HCO<sub>3</sub>. It is evident from the data that five principal components (PCs) account for 98.993% of the total variability in the data set (**Table 3**). The loadings are the

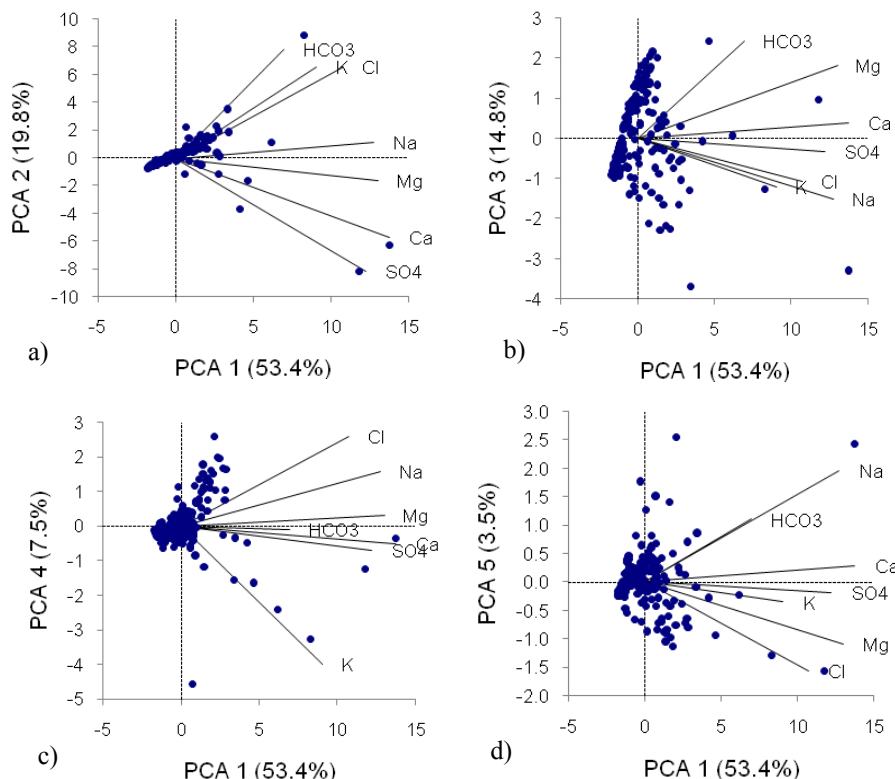


Figure 9. Loading plots of component 2 to 5 on component 1 of the groundwater quality (a: shallow zone aquifer in the recharge area made of shale, quartzites b) dolomitic aquifer, c) acidic crystalline rocks aquifer such as granite and gneisses, d) weathered zone aquifer in low altitude areas).

Table 3. Weights of the five principal components.

Variable	PC_1	PC_2	PC_3	PC_4	PC_5
Ca	0.878	-0.426	0.102	-0.071	0.047
Mg	0.831	-0.119	0.489	0.045	-0.181
Na	0.814	0.085	-0.408	0.225	0.325
K	0.577	0.488	-0.323	-0.565	-0.058
Cl	0.687	0.489	-0.287	0.367	-0.260
SO <sub>4</sub>	0.781	-0.605	-0.089	-0.099	-0.032
HCO <sub>3</sub>	0.444	0.583	0.653	-0.014	0.183
eigenvalues	3.737	1.387	1.037	0.523	0.246
Cum. %	53.387	73.200	88.019	95.485	98.993

correlations between the original variable and the individual components. The values indicate roughly the relative contribution/importance of each variable in each principal component [31]. The result shows that Ca, Mg, Na and SO<sub>4</sub> have maximum loading on PC<sub>1</sub>, which could be considered as a principal chemical constituent of the aquifer in the area except SO<sub>4</sub>, which is derived from AMD, while HCO<sub>3</sub> has highest loading on PC<sub>2</sub> and PC<sub>3</sub>.

The clustering of most of the data points to the origin

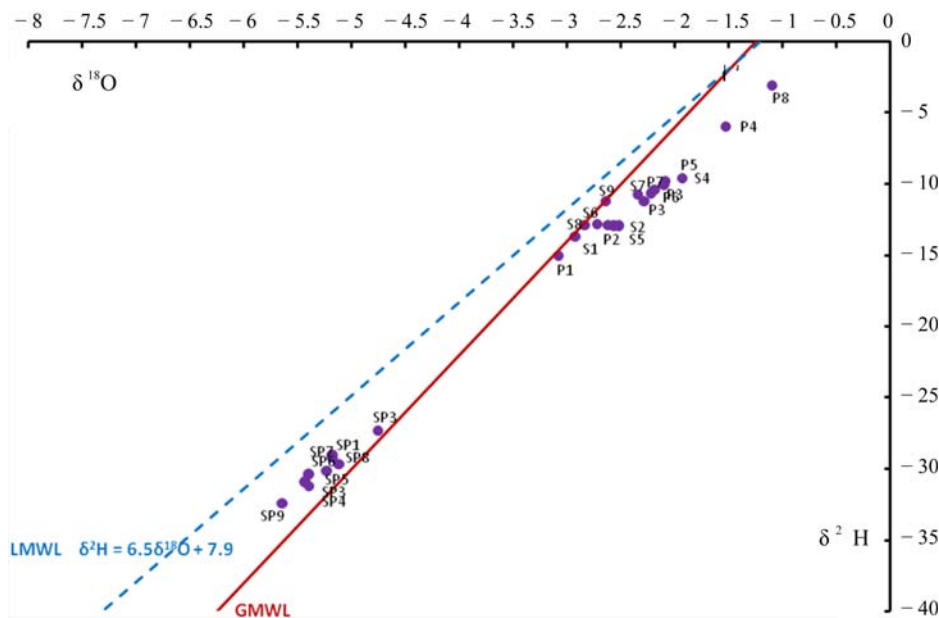
shows significantly dissimilar data. Figure 9(a) shows loading of PC<sub>2</sub> on PC<sub>1</sub> with HCO<sub>3</sub>, K and Cl. Possible occurrence could be shallow groundwater with a source from a shale dominated recharge zone, which is the main lithology on the northern slope of the watershed. Figure 9(b) shows loading of PC<sub>3</sub> on PC<sub>1</sub> with HCO<sub>3</sub>, Mg and Ca. Possible source is likely to be dolomitic rocks. Figure 9(c) shows loading of PC<sub>4</sub> on PC<sub>1</sub> with Cl, Na and Mg. For such type of source rock is deduced to be an aquifer of acidic crystalline rock such as granite and gneiss. However, Cl can be generated from marine precipitation, urban pollution etc. Figure 9(d) shows loading of PC<sub>5</sub> on PC<sub>1</sub> with Na and HCO<sub>3</sub>, which represents particularly shallow circulating groundwater within the weathered zones rich in clay.

### 5.3. Environmental Isotope Data

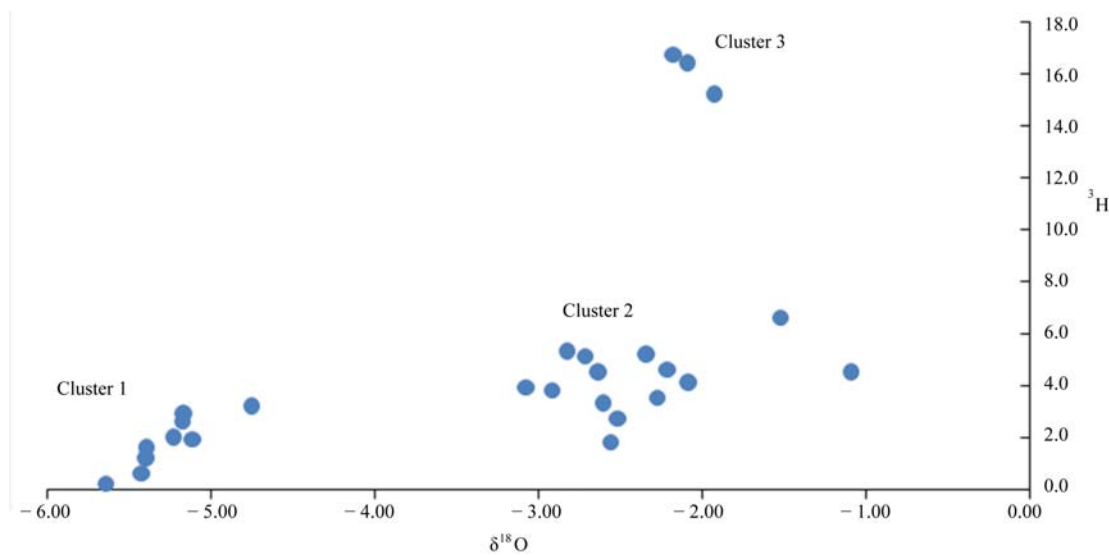
Environmental isotopes such as δ<sup>18</sup>O, δ<sup>2</sup>H and <sup>3</sup>H are widely used to gain some insight into the subsurface flow and recharge condition. These isotopes have been used to study the recharge condition in the sedimentary sequence in South Africa [32,33] to characterize the recharge mechanism and geochemical processes in the arid environment. The plots on Figure 10 indicate that

**Table 4. Environmental isotopes in the Johannesburg area.**

Sample code	E (UTM)	N (UTM)	T (°C)	pH	TDS (mg/L)	EC ( $\mu$ S/cm)	Eh (mv)	$\delta^{18}\text{O}$ (‰)	$\delta^2\text{H}$ (‰)	Tritium (T.U.)
SP1	7134258	579896	22.25	7.51	116.2	181.4	-15.7	-5.17	-29.0	2.9 $\pm$ 0.3
SP2	7134851	579772	22.45	8.47	120.7	188.9	-64.5	-4.75	-27.3	3.2 $\pm$ 0.3
SP3	7137898	578656	22.20	7.65	140.2	220.0	-22.2	-5.42	-31.0	0.6 $\pm$ 0.2
SP4	7137903	578704	21.75	7.77	145.5	228.0	-29.5	-5.40	-31.2	1.2 $\pm$ 0.2
SP5	7137958	578194	21.65	8.53	110.6	172.6	-73.2	-5.39	-30.4	1.6 $\pm$ 0.3
SP6	7137953	578184	20.95	8.59	87.4	136.5	-75.2	-5.23	-30.1	2.0 $\pm$ 0.3
SP7	7137893	578150	21.20	8.52	77.8	121.4	-76.5	-5.11	-29.6	1.9 $\pm$ 0.3
SP8	7137885	578140	23.00	8.38	108.1	168.9	-63.1	-5.17	-29.2	2.6 $\pm$ 0.3
SP9	7137818	578117	21.40	7.66	110.0	172.3	-31	-5.64	-32.4	0.2 $\pm$ 0.2
S1	574686	7114806	20.20	7.52	118.8	185.3	-0.1	-2.92	-13.7	3.8 $\pm$ 0.3
S2	572242	7112068	20.20	3.55	1502.0	2350.0	198.1	-2.56	-13	1.8 $\pm$ 0.3
S3	572107	7119673	22.10	7.53	216.0	337.0	-18.2	-2.22	-10.7	4.6 $\pm$ 0.4
S4	571951	7121518	21.30	7.57	217.0	339.0	-21.9	-2.09	-9.8	4.1 $\pm$ 0.4
S5	577604	7124265	20.60	7.82	225.0	352.0	-35.3	-2.52	-13	2.7 $\pm$ 0.3
S6	577604	7124265	20.60	7.90	194.3	303.0	-38.1	-2.61	-12.9	3.3 $\pm$ 0.3
S7	574187	7123145	20.50	7.82	229.0	358.0	-25.4	-2.28	-11.2	3.5 $\pm$ 0.3
S8	598990	7106948	21.10	8.12	135.8	212.0	-48.9	-2.83	-12.9	5.3 $\pm$ 0.4
S9	600335	7107863	21.40	8.24	112.6	176.0	-39.8	-2.64	-11.2	4.5 $\pm$ 0.4
P1	7103843	598632	15.40	7.44	171.9	269.0	-32.6	-3.08	-15.0	3.9 $\pm$ 0.3
P2	7106029	599933	13.20	7.20	106.0	167.0	-15	-2.72	-12.8	5.1 $\pm$ 0.4
P3	7111252	601923	11.20	7.99	149.1	233.0	-40.5	-2.35	-10.7	5.2 $\pm$ 0.4
P4	7146708	589660	16.80	8.21	181.4	284.0	-52	-1.53	-6.0	6.6 $\pm$ 0.4
P5	7150413	580760	20.40	9.60	146.4	229.0	-125.1	-1.93	-9.6	15.2 $\pm$ 0.7
P6	7155277	584190	21.20	9.37	148.9	232.0	-111.5	-2.09	-10.1	16.4 $\pm$ 0.7
P7	7160697	579527	20.40	8.84	158.1	247.0	-87.7	-2.18	-10.4	16.7 $\pm$ 0.7
P8	7132300	622947	18.70	7.88	210.0	328.0	-39.9	-1.10	-3.1	4.5 $\pm$ 0.4



(a)



(b)

**Figure 10. Environmental isotope distribution in the study area.**

there are different aquifers in the area with variable recharge history. All data (Table 4) falling below the Local Meteoric Water Line (LMWL) indicate recharge right after evaporation. One of the highly productive springs in the Sterkfontein area within the Cradle of Human Kind World Heritage Site, identified as Ngosi spring, emerges as a group of springs from dissolution cavities with a discharge of about 145 L/s. For this particular spring the measured  $\delta^{18}\text{O}$ ,  $\delta^2\text{H}$  and  $^3\text{H}$  values are  $-5.42\text{‰}$ ,  $-31.0\text{‰}$ ,  $0.6 \pm 0.2$  TU respectively. The data show that the

spring contains an old water with more than five decades in circulation (low  $^3\text{H}$ ) and high altitude recharge or deep circulation (depleted stable isotopes) indicating the presence of some deeper circulation chamber in the dolomitic aquifer.

Stable isotope plot with respect to meteoric water lines (Figure 10(a)) portrays two clusters with deep circulation (Cluster 1) within the dolomitic aquifer, slower moving older component of groundwater (depleted  $\delta^{18}\text{O}$  and low  $^3\text{H}$ ), and the shallow circulation within the

weathered crystalline rocks besides streams recharge by interflows (cluster 2). The variation diagrams are found to be supportive of the occurrence of a dominantly shallow groundwater circulation. **Figure 8(b)** reveals three clusters with the third cluster (cluster 3) most likely originating from industrial waste input of the South African Nuclear Energy Corporation whereas the remaining two clusters conform to the shallow circulation.

## 6. Conclusions

The extreme precarious geoenvironmental setting of the greater Johannesburg area highly impacted the groundwater resource where the influence of the lithology and pollution on the water composition is manifested through the variation in the groundwater quality. The study revealed the loss of major ions with depth in the crystalline aquifers of the study area showing the main movement mechanism of chemical constituents is diffusive type, which suggests the decrease in water velocity with depth. It is highly likely that diffusive circulation is caused by fracture sealing by calcite precipitation that reduces the available permeability with depth. The enhanced dissolution of minerals due to acid mine drainage has important contribution for the evolution of ions in the groundwater. Rapid infiltration through fractures or dissolution cavities results in low degree of water-rock interaction and hence, low salinity waters exist within the dolomitic aquifer with low chloride and depleted  $\delta^{18}\text{O}$ , indicating recharge from relatively high altitude. It was observed that slow moving older groundwater (depleted  $\delta^{18}\text{O}$  and low  $^3\text{H}$ ) occurs within the weathered crystalline aquifer and it mixes with relatively young and shallow moving locally recharged water.

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